



On secondary new particle formation in China

Journal:	<i>Frontiers of Environmental Science and Engineering</i>
Manuscript ID	FESE-2016-0074.R1
Manuscript Type:	Special Issue: Progresses in understanding secondary air pollution
Date Submitted by the Author:	n/a
Complete List of Authors:	<p>Kulmala, Markku; University of Helsinki Petäjä, Tuukka; University of Helsinki; Nanjing University and University of Helsinki Kerminen, Veli-Matti; University of Helsinki Kujansuu, Joni; University of Helsinki Ruuskanen, Taina; University of Helsinki Ding, Aijun; Institute for Climate and Global Change Research; Nanjing University and University of Helsinki Nie, Wei; University of Helsinki; Nanjing University and University of Helsinki ; Institute for Climate and Global Change Research HU, Min; State Key Joint Laboratory of Environmental Simulation and Pollution Control, College of Environmental Sciences and Engineering Wang, Zhibin; Max Planck Institute for Chemistry; State Key Joint Laboratory of Environmental Simulation and Pollution Control, College of Environmental Sciences and Engineering Wu, Zhijun; State Key Joint Laboratory of Environmental Simulation and Pollution Control, College of Environmental Sciences and Engineering Wang, Lin; Fudan University, Worsnop, Douglas; University of Helsinki; Aerodyne Research Inc</p>
Keywords:	aerosol particles, heavily-polluted environments, condensation sink, new particle production, megacities
Speciality:	Formation < POLLUTION, Particle-phase species < POLLUTION, Mass spectrometry < ANALYTICAL METHODS

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

For Review Only

1
2
3 1 **The original question from guest editor: your view or comments or directions**
4
5 2 **related to secondary aerosol formation in China**
6
7

8
9 3

10
11 4 *Journal: Front. Environ. Sci. Eng*
12
13

14
15 5

16
17 6 Title:
18
19

20 7 **On secondary new particle formation in China**
21
22

23
24 8

25
26 9 Running title:
27
28

29 10 **On secondary new particle formation in China**
30
31

32
33 11

34
35 12 Correspondence author
36
37

38 13 Markku Kulmala
39
40

41 14 University of Helsinki
42
43

44 15 P.O. Box 64
45
46

47 16 Helsinki, FI 00014, Finland
48
49

50 17 markku.kulmala@helsinki.fi
51
52

53 18 Tel: +358-40-5962311
54
55

1
2
3 19 Fax: +358-9-19150717
4
5
6
7

8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

21 Markku Kulmala¹

22 ¹University of Helsinki

23 P.O. Box 64

24 Helsinki, FI 00014, Finland

25 markku.kulmala@helsinki.fi

26

27 Tuukka Petäjä^{1,2}

28 ¹University of Helsinki

29 P.O. Box 64

30 Helsinki, FI 00014, Finland

31 tuukka.petaja@helsinki.fi

32 ²Joint International Research Laboratory of Atmospheric and Earth System Sciences

33 (JirLATEST)

34 Nanjing University and University of Helsinki

35 22 Hankou Road, Nanjing, CN 210093, China

36

37 Veli-Matti Kerminen¹

1
2
3 38 ¹University of Helsinki
4
5

6 39 P.O. Box 64
7
8

9 40 Helsinki, FI 00014, Finland
10
11

12 41 veli-matti.kerminen@helsinki.fi
13
14

15 42
16
17

18 43 Joni Kujansuu¹
19
20

21 44 ¹University of Helsinki
22
23

24 45 P.O. Box 64
25
26

27 46 Helsinki, FI 00014, Finland
28
29

30 47 joni.kujansuu@helsinki.fi
31
32

33 48
34
35

36 49 Taina Ruuskanen¹
37
38

39 50 ¹University of Helsinki
40
41

42 51 P.O. Box 64
43
44

45 52 Helsinki, FI 00014, Finland
46
47

48 53 taina.ruuskanen@helsinki.fi
49
50

51 54
52
53

54 55 Aijun Ding^{2,3}
56
57

1
2
3 56 ²Joint International Research Laboratory of Atmospheric and Earth System Sciences
4
5 57 (JirLATEST)

6
7
8 58 Nanjing University and University of Helsinki

9
10
11 59 22 Hankou Road, Nanjing, CN 210093, China

12
13
14 60 ³Institute for Climate and Global Change Research

15
16
17 61 Nanjing University

18
19
20 62 Nanjing, Jiangsu, CN 210000, China

21
22
23 63 dingaj@nju.edu.cn

24
25
26 64

27
28
29 65 Wei Nie^{1,2,3}

30
31
32 66 ¹University of Helsinki

33
34
35 67 P.O. Box 64

36
37
38 68 Helsinki, FI 00014, Finland

39
40
41 69 ²Joint International Research Laboratory of Atmospheric and Earth System Sciences
42
43 70 (JirLATEST)

44
45
46 71 Nanjing University and University of Helsinki

47
48
49 72 22 Hankou Road, Nanjing, CN 210093, China

50
51
52 73 ³Institute for Climate and Global Change Research

53
54
55 74 Nanjing University

1
2
3 75 Nanjing, Jiangsu, CN 210000, China
4
5

6 76 niewei@nju.edu.cn
7
8

9 77
10

11 78 Min Hu⁴
12
13

14
15 79 ⁴State Key Joint Laboratory of Environmental Simulation and Pollution Control,
16

17 80 College of Environmental Sciences and Engineering, Peking University
18

19 81 Beijing, CN 100871, China
20
21

22 82 minhu@pku.edu.cn
23
24

25 83
26
27

28 84 Zhibin Wang^{4,5}
29
30

31 85 ⁴State Key Joint Laboratory of Environmental Simulation and Pollution Control,
32

33 86 College of Environmental Sciences and Engineering, Peking University
34

35 87 Beijing, CN 100871, China
36
37

38 88 ⁵Max Planck Institute for Chemistry
39
40

41 89 Hahn-Meitner-Weg 1, Mainz, DE 55128, Germany
42
43

44 90 zhibin.wang@mpic.de
45
46

47 91
48
49

50 92 Zhijun Wu⁴
51
52
53
54
55
56
57
58
59
60

1
2
3 93 ⁴State Key Joint Laboratory of Environmental Simulation and Pollution Control,
4
5 94 College of Environmental Sciences and Engineering, Peking University
6
7 95 Beijing, CN 100871, China
8
9

10 96 zhijunwu@pku.edu.cn
11
12
13
14 97

15
16 98 Lin Wang⁶
17

18
19 ⁶Fudan University, Department of Environmental Science & Engineering
20
21

22 100 Shanghai, CN 200433, China
23
24

25 101 lin_wang@fudan.edu.cn
26
27

28 102
29
30

31 103 Douglas R. Worsnop^{1,7}
32
33

34 104 ¹University of Helsinki
35
36

37 105 P.O. Box 64
38
39

40 106 Helsinki, FI 00014, Finland
41
42

43 107 ⁷Aerodyne Research Inc
44
45

46 108 Billerica, MA, USA 01821
47
48

49 109 worsnop@aerodyne.com
50
51

52 110
53
54

55
56 111 **Abstract**
57

1
2
3 112
4
5

6 113 Formation of new atmospheric aerosol particles is a global phenomenon that has been
7
8 114 observed to take place in even heavily-polluted environments. However, in all
9
10 115 environments there appears to be a threshold value of the condensation sink (due to
11
12 116 pre-existing aerosol particles) after which the formation rate of 3 nm particles is no
13
14 117 longer detected. In China, new particle production has been observed at very high
15
16 118 pollution levels (condensation sink about 0.1 s^{-1}) in several megacities, including
17
18 119 Beijing, Shanghai and Nanjing as well as in Pearl River Delta (PRD). Here we
19
20 120 summarize the recent findings obtained from these studies and discuss the various
21
22 121 implications these findings will have on future research and policy.
23
24
25
26
27 122
28
29

30 123 **1. Background**

31
32
33 124
34
35

36 125 Atmospheric aerosol particles affect our life and its quality in multiple ways. First of
37
38 126 all, the interaction between aerosols and climate system is the dominant uncertainty in
39
40 127 predicting the radiative forcing and future climate [1]. Secondly, aerosol particles
41
42 128 deteriorate both human health and visibility, especially in urban areas [2, 3]. Thirdly,
43
44 129 aerosol particles modify the intensity and distribution of radiation that reaches the
45
46 130 Earth's surface, having direct influences on photosynthesis and terrestrial carbon sink
47
48 131 [4]. Better understanding of the various effects in the atmosphere requires detailed
49
50 132 information on how different sources (including those related to the biosphere) and
51
52 133 atmospheric transformation processes modify the properties of aerosol particle
53
54 134 populations.
55
56
57
58
59
60

1
2
3 135
4
5

6 136 One of the most important phenomena associated with the atmospheric aerosol
7
8 137 number concentrations is the secondary formation of new aerosol particles. This
9
10 138 includes the production of molecular clusters from gaseous precursor vapors, the
11
12 139 activation and growth of some of these clusters to detectable sizes, and the further
13
14 140 growth up to the sizes at which the particles may act as cloud condensation nuclei
15
16
17 141 [e.g. 5, 6]. Although atmospheric new particle formation has been observed to take
18
19 142 place almost everywhere at favorable conditions in the boundary layer [7], our
20
21 143 knowledge about this phenomenon is still far from perfect [5, 8]. The current
22
23 144 knowledge gaps in this regard range from the basic process-level understanding of
24
25 145 secondary atmospheric aerosol formation to its connection with anthropogenic
26
27 146 activities, biogenic emissions, atmospheric chemistry, and ultimately with climate
28
29 147 change and human health.
30
31

32
33 148
34
35

36 149 Secondary formation of new atmospheric aerosol particles is typically initiated by
37
38 150 photochemical reactions in the gas phase, so that especially the production of
39
40 151 extremely low volatility vapors like sulfuric acid [9, 10, 11] and highly-oxidized
41
42 152 organic compounds [e.g. 12, 13, 14] is crucial. Pre-existing aerosol particles act as a
43
44 153 sink for the low-volatile vapors, as well as for small clusters and growing
45
46 154 nanoparticles, thereby hindering or even suppressing atmospheric new particle
47
48 155 formation [e.g. 15, 16, 17]. The atmospheric new particle formation is affected by
49
50 156 several meteorological quantities and phenomena, particularly in the planetary
51
52 157 boundary layer, including the intensity of solar radiation and atmospheric mixing
53
54
55
56
57
58
59
60

1
2
3 158 processes. The recent findings indicate that critical clusters may be surprisingly small
4
5 159 in size, if existing at all, under atmospheric conditions [e.g. 18], and thus treatable by
6
7 160 advanced quantum chemistry methods [19]. It is very probable that the atmospheric
8
9 161 new particle formation is a two-step process, i.e. initial clustering and then
10
11 162 condensational growth after activation of clusters, as suggested by Kulmala et al. [21]
12
13 163 and verified by Kulmala et al. [18]. A summary of the current understanding of gas-
14
15 164 to-particle conversion is presented by Kulmala et al. [5].

16
17
18
19 165 New aerosol particles formed in the atmosphere become climatically important when
20
21 166 they reach sizes larger than about 50–100 nm in diameter [6]. Particles of this size
22
23 167 and larger are able to act as cloud condensation nuclei and scatter visible light,
24
25 168 thereby affecting cloud microphysical properties [e.g. 22], reducing the fraction of
26
27 169 solar radiation reaching the Earth's surface and contributing to visibility degradation
28
29 170 [e.g. 23]. Furthermore, health effects of airborne particles are related not only to the
30
31 171 amount and toxicity of the particulate material, but also to the particle size because
32
33 172 this property has a large effect on whether or not a particle is able to penetrate into the
34
35 173 lungs [e.g. 20] and even further into the blood circulation [e.g. 24].

36
37
38
39
40
41 174

42
43 175 The rapid, large-scale urbanization and industrialization of China are unique in
44
45 176 history. Consequently, China's air pollution situation has worsened dramatically
46
47 177 during the last 2–3 decades as emissions from industry, energy production and traffic
48
49 178 have increased. China is currently responsible for 30–35 % of the global SO₂, NO_x,
50
51 179 CO and Particulate mass (PM) emissions and 40% of global particle number (PN)
52
53 180 emission in the 20–1000 nm size range (see <http://gains.iiasa.ac.at/gains3/>).

1
2
3 181 Atmospheric concentrations of primary and secondary pollutants in China are 10 to
4
5 182 100 times (sometimes even 1000 times) higher than currently in Europe or Northern
6
7 183 America. However, highly non-linear processes, such as atmospheric chemistry and
8
9 184 aerosol dynamics, transform the urban pollution cocktail and generate secondary
10
11 185 pollution, such as ultrafine particles and ozone, during their residence in the
12
13 186 atmosphere [25, 26]. The fact that new particle formation does occur in polluted
14
15 187 Chinese megacities like Beijing [27] and Shanghai [28], or even during dust-storms
16
17 188 [29, 30], suggests that there are several major physical and chemical mechanisms in a
18
19 189 heavily-polluted atmosphere that have not been recognized before and may not even
20
21 190 be operating in clean or moderately-polluted environments. At present, atmospheric
22
23 191 air pollution in China threatens the health of hundreds of millions of people [e.g. 3,
24
25 192 31], and causes major problems to the environment and economy as a whole by
26
27 193 decreasing, e.g. severely the agricultural and industrial productivity of the nation as a
28
29 194 whole. This pollution also reduces visibility, thereby decreasing the attraction of these
30
31 195 mega-cities for tourists, and hinders the possibilities to use solar energy a source for a
32
33 196 clean energy on a local scale.
34
35
36
37
38
39
40
41

42 198 A holistic scientific understanding on the atmospheric phenomena associated with air
43
44 199 quality as a whole, as well as on the connection between air quality and climate, is
45
46 200 lacking at the moment [31-33]. Together with emission reductions, the key way to get
47
48 201 forward is to perform long-term, continuous and comprehensive observations on
49
50 202 aerosol particles (mass, number, chemical composition, optical properties), on
51
52 203 concentrations of trace gases (SO₂, NO_x, CO, VOCs, sulphuric acid, HONO, HNO₃,
53
54 204 NH₃ etc.), and on atmospheric oxidant levels (O₃, HO_x, RO_x, NO₃, Criegee
55
56
57
58
59
60

1
2
3 205 intermediates etc.), as well as on greenhouse gas concentrations [31]. With a network
4
5 206 of such observation stations [34], we will be able to understand the interactions and
6
7 207 feedbacks associated with the urban pollution mixture [e.g. 35-37], and ultimately, be
8
9 208 ready to make targeted strategies for the pollution control. In the following we take
10
11 209 recent advances in studying secondary new aerosol formation in China as an example
12
13 210 to show how increased process-level understanding will help us to understand air
14
15 211 quality-climate-weather interactions and how the feedbacks and interactions affect the
16
17 212 air quality in highly-polluted environments such as those frequently encountered in
18
19 213 Chinese megacities.
20
21
22
23
24
25
26

27 215 **2. Results from recent studies on New Particle Formation in China**

28
29
30
31
32

33 217 New particle formation events have been observed in many different locations in
34
35 218 China, including coastal/marine, rural, regional and polluted urban environments [28,
36
37 219 30, 38-47].
38
39
40
41
42

43 221 The first long-term study on NPF events was performed in the urban of Beijing at
44
45 222 PKU Urban Atmosphere Environment MonitoRing Station (PKUERS), starting at
46
47 223 2004 [27, 48, 49]. On average, every fifth day (~21%) displayed a NPF event [50]. An
48
49 224 evident seasonal variation profile for NPF events was observed, showing that a high
50
51 225 frequency the NPF events (~ 40%) occurred during the spring and winter [27, 50],
52
53 226 while fewer events were observed in summer [51, 52]. The observed formation rates
54
55
56
57
58
59
60

1
2
3 227 of 3-nm particles and their growth rates were in the ranges of $3.3\text{-}81.4\text{ cm}^{-3}\text{ s}^{-1}$ and
4
5 228 $0.1\text{-}11.2\text{ nm h}^{-1}$ [27, 50, 53], respectively.
6
7

8
9 229

10
11 230 Generally, NPF is an unexpected phenomenon in the polluted atmosphere of China
12
13 231 due to typically high loadings of pre-existing aerosol particles. For example, the mean
14
15 232 condensation sink (CS, [54]) values during the nucleation event days were 0.025 s^{-1}
16
17 233 ($0.003\text{-}0.086\text{ s}^{-1}$) and 0.026 s^{-1} ($0.004\text{-}0.082\text{ s}^{-1}$) at the rural (Kaiping) and urban
18
19 234 (Beijing) environments, respectively, which are approximately 5 to 10 times higher
20
21 235 than the values of CS observed in clean environments [55-57]. This high
22
23 236 concentration of pre-existing aerosol particles significantly inhibits the growth of
24
25 237 newly-formed particles. In fact, the observed NPF event is an end product of the
26
27 238 competition between the low-volatile vapor sources (such as SO_2 or sulfuric acid) and
28
29 239 sinks (such as pre-existing particles), as shown by Kulmala et al. [55]. The abundant
30
31 240 SO_2 emissions and high oxidation capacity in the polluted atmosphere of China
32
33 241 indicate that there is a sufficient source of sulfuric acid [40, 52]. Therefore, in the case
34
35 242 of both higher source and sink, their inter-competition is the most likely factor that
36
37 243 determines the occurrences of NPF events in polluted environments.
38
39
40
41
42

43
44 244

45
46 245 Two years (2011-2013) of continuous particle number size distribution measurements
47
48 246 were conducted at the Station for Observing Regional Processes of the Earth System
49
50 247 (SORPES [35, 36]) station about 20 km northeast of urban Nanjing. The location can
51
52 248 be considered as a regional background site of Yangtze River Delta in eastern China.
53
54 249 During this time period, 44% of the sampling days were NPF event days (see Figure 1
55
56

1
2
3 250 as an example). The formation rates of 6-nm particles varied from 0.24 to $10.9 \text{ cm}^{-3} \text{ s}^{-1}$
4
5 251 $^{-1}$, the subsequent particle growth rates varied from 3.6 to 23 nm h^{-1} , and the values of
6
7 252 CS during the event days varied from 0.007 to 0.068 s^{-1} [47]. Most of the NPF events
8
9
10 253 took place in spring, summer and autumn with the frequencies of 55, 54 and 49 %,
11
12 254 respectively, whereas only 15 events (11.2%) were observed in winter.

13
14
15 255

16
17
18 256 **Figure 1.** A typical nucleation event measured using Air Ion Spectrometer (AIS) at
19
20 257 the SORPES station, Nanjing, in China. The background cluster ions are seen in both
21
22 258 negative and positive ion modes in the sub-2 nm size range. Negative ion clusters are
23
24 259 smaller than positive ones. The new particle formation is seen in both polarities
25
26 260 starting at around 8.30 am. Here J_6 is $1.8 \text{ cm}^{-3} \text{ s}^{-1}$ and GR (6-30 nm) is 6.6 nm h^{-1} .

27
28
29
30 261

31
32
33 262 The typical NPF event in Nanjing is shown in Figure 1. In Nanjing, many of the NPF
34
35 263 events occurred on the days associated with heavy pollution. As shown by Xie et al.
36
37 264 [30], frequent NPF events were observed when the $\text{PM}_{2.5}$ and PM_{10} concentrations
38
39 265 were in excess of $100 \text{ } \mu\text{g m}^{-3}$ and $200 \text{ } \mu\text{g m}^{-3}$, respectively. The reason for this is still
40
41 266 an open question. One hypothesis is that nucleation can be promoted by
42
43 267 heterogeneous reactions on the surface of the dust [29, 30]. This is supported by many
44
45 268 observations from both SORPES station and another mountain top site, Mt. Heng in
46
47 269 southern China. In the spring of 2009, relatively high new-particle formation rates
48
49 270 ($0.46 \text{ cm}^{-3} \text{ s}^{-1}$) and growth rates (7.2 nm h^{-1}) were observed when the loading of pre-
50
51 271 exist particles was higher than $600 \text{ } \mu\text{g m}^{-3}$ at Mt. Heng. Combined with laboratory
52
53
54
55
56
57
58
59
60

1
2
3 272 investigations [58], dust-induced heterogeneous photochemical processes were
4
5 273 supposed to provide additional gaseous oxidants to promote the NPF [29].
6
7
8
9 274

10
11 275 In urban Shanghai, particle size distributions were measured from November 2013 to
12
13 276 January 2014 on the rooftop of a teaching building (31°18'N, 121°30'E) on the
14
15 277 campus of Fudan University [28], which can be regarded as an urban site. During this
16
17 278 62-day campaign, 13 NPF events were identified with strong bursts of sub-3 nm
18
19 279 particles and subsequent fast growth of these particles. The observed nucleation rate
20
21 280 ($J_{1.34}$), formation rate of 3 nm particles (J_3), and CS were in the ranges of 112.4-
22
23 281 $271.0 \text{ cm}^{-3} \text{ s}^{-1}$, $2.3\text{-}19.2 \text{ cm}^{-3} \text{ s}^{-1}$ and $0.030\text{-}0.10 \text{ s}^{-1}$, respectively. The growth rates of
24
25 282 the formed clusters and nanoparticle showed a clear size dependence, with average
26
27 283 values of $\text{GR}_{1.35\sim 1.39}$, $\text{GR}_{1.39\sim 1.46}$, $\text{GR}_{1.46\sim 1.70}$, $\text{GR}_{1.70\sim 2.39}$, $\text{GR}_{2.39\sim 7}$ and
28
29 284 $\text{GR}_{7\sim 20}$ being 1.6 ± 1.0 , 1.4 ± 2.2 , 7.2 ± 7.1 , 9.0 ± 11.4 , 10.9 ± 9.8 and $11.4\pm 9.7 \text{ nm h}^{-1}$,
30
31 285 respectively. Nucleation of particles during this campaign might be explained by the
32
33 286 activation theory, since the formation rate of the smallest particles was proportional to
34
35 287 a 0.65 ± 0.28 power of the sulfuric acid proxy. In addition, ammonia was very likely
36
37 288 associated with NPF events, as the new particle formation rate was positively
38
39 289 correlated with the concentration of gas-phase ammonia. The estimated sulfuric acid
40
41 290 concentration was sufficient to explain the growth of 1.34–3 nm particles, but its
42
43 291 contribution became smaller as the particle grew in size.
44
45
46
47
48
49
50
51 292

1
2
3 293 The observed new particle rates, condensation sink and particle growth rates in the
4
5 294 three megacities, i.e. Beijing, Nanjing and Shanghai, are of the same order of
6
7 295 magnitude. These similarities reflect the urban nature of the Beijing and Shanghai
8
9 296 sites, and hint that the Nanjing site, although considered as a regional background site
10
11 297 of Yangtze River Delta in eastern China, might be characterized with a similar
12
13 298 competition between the sources and sinks of low-volatility vapors. The seasonal
14
15 299 pattern of the NPF frequency is very different between the two sites having long-term
16
17 300 measurements, Beijing and Nanjing, in addition which the annual-averaged NPF
18
19 301 frequency is clearly higher in Nanjing. The fundamental reason for these differences
20
21 302 lies probably in a delicate balance between the factors that favor or suppress new
22
23 303 particle formation and growth. At both Beijing and Nanjing, for example, NPF is
24
25 304 favored by a low ambient relative humidity and low CS, whereas no consistent pattern
26
27 305 can be seen between the occurrence of NPF and either the ambient temperature or
28
29 306 sulfur dioxide concentration [27, 47]. The fact that high values of CS tend to suppress
30
31 307 NPF is fully in line with theoretical expectations [16, 17], and it might explain the low
32
33 308 NPF frequency observed in Shanghai during polluted winter conditions [28]. There is
34
35 309 strong, yet indirect evidence that NPF events in these three megacities are connected
36
37 310 to sulfuric acid [28, 29, 40, 52]. However, it is premature to conclude that the exact
38
39 311 nucleation mechanisms are identical in three megacities without direct measurements
40
41 312 of chemical composition of nucleating clusters and ions.
42
43
44
45
46
47
48
49
50
51

52 314 Besides the direct connection between pre-exist aerosols (e.g. mixed dust) and NPF, a
53
54 315 recent study found that biomass burning particles can enhance the conversion rate of
55
56
57
58
59
60

1
2
3 316 NO₂ to HONO which is one of the main sources of OH and can in turn promote the
4
5 317 formation of secondary aerosol mass and number [59]. Furthermore, it was found that
6
7 318 when biomass burning particles are mixed with anthropogenic pollution, the HONO
8
9 319 production potential from the conversion of NO₂ to HONO tend to be enhanced even
10
11 320 more. Given that biomass burning particles are easily mixed with anthropogenic
12
13 321 pollution in eastern China, their influences on the HONO budget, radical pool, and
14
15 322 thus the formation of secondary aerosols are expected to be important [59].
16
17
18
19
20
21
22

23 324 Heterogeneous, or multi-phase, processes influence the secondary aerosol formation.
24
25 325 For example, most of aerosol sulfate has been believed to be formed from
26
27 326 heterogeneous or aqueous-phase processes (cloud processes). Ozone and hydrogen
28
29 327 peroxide are the major oxidants to drive these processes. Recent studies have shown
30
31 328 that NO₂ can also be an important oxidant to convert SO₂ to sulfate when mineral dust
32
33 329 and biomass burning plumes are present [30, 60]. Especially during the biomass
34
35 330 burning-induced haze events [30], the oxidation processes by NO₂ became critical
36
37 331 when the formation of other oxidants were suppressed. More interestingly, one of the
38
39 332 “by-products” of the reaction of SO₂ and NO₂ is HONO, which can further enhance
40
41 333 the atmospheric oxidation capacity. All these observations suggest that our current
42
43 334 understanding on secondary aerosol formation processes need to be revised.
44
45
46
47
48
49
50
51

52 336 **3. On future NPF studies**
53
54
55 337

1
2
3 338 The importance of secondary aerosols has become apparent during the last decades,
4
5 339 so there is an increasing need for understanding their formation mechanisms and
6
7 340 atmospheric dynamics in detail. Although several field campaigns and a few long-
8
9 341 term (over several years) observations on NPF have already been conducted in China,
10
11 342 we need to perform additional long-term measurements, preferable continuous and
12
13 343 comprehensive observations utilizing the full capacity of the current state-of-the-art
14
15 344 instruments.
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

345
346 In the coming decade, we need to utilize the full capacity of new aerosol and ion
347 instruments, such as the Particle Size Magnifier (PSM, [61]), Neutral cluster and Air
348 Ion Spectrometer (NAIS, [62]) and Sigma [63]. With these instruments, we will be
349 able to detect and analyze the frequency of NPF events, as well as to determine cluster
350 concentrations, particle formation rates and size-dependent particle growth rates [e.g.
351 64, 65]. Furthermore, we will be able to quantify the contribution of ion and neutral
352 pathways to NPF [66, 67].
353

354 The aerosol and ion instruments together with the high-resolution mass spectrometers,
355 such as Atmospheric Pressure interface – Time of Flight mass spectrometer (APiTOF,
356 [68] and Chemical Ionization APiTOF [69], will make it possible to connect the NPF
357 to the concentrations of different vapors participating in this process. Such vapors
358 include sulfuric acid [9, 70, 71], ammonia [72], amines [12, 73] and organic vapors
359 [11, 13, 74]. Furthermore, the mass spectrometers can also be utilized in determining

1
2
3 360 atmospheric radical concentrations [75, 76, 77] responsible for the oxidation of
4
5 361 precursor vapors in the atmosphere.
6
7

8 362
9

10
11 363 To support the NPF analysis, aerosol number size distributions need to be measured
12
13 364 with harmonized instruments [78], enabling quantification of the condensation sink of
14
15 365 a pre-existing particle population. On-line chemical analysis is important as well,
16
17 366 since such information can be used to attributing the relative contributions of different
18
19 367 aerosol sources [e.g. 26, 79].
20
21
22

23 368
24
25

26 369 In order to have reliable data which can also be compared from one site to another,
27
28 370 instruments need to be calibrated often enough in the laboratory. This should be
29
30 371 conducted within specific calibration centers. In order to assure the data quality, open
31
32 372 data flows and joint data analysis are preferable, which will lead to joint publications
33
34 373 and provides novel avenues to exploit the data to improve both regional air quality
35
36 374 and global climate.
37
38

39
40 375
41
42

43 376 **4. Capacity building** 44

45 377
46
47

48 378 Capacity building related to scientists, engineers and technicians operating
49
50 379 instruments and stations are necessary pre-requisites for obtaining good data. For
51
52 380 example, a proper use of instruments will optimize the efforts, improve the data
53
54 381 quality and enhance data and publication flows.
55
56

1
2
3 382
4
5
6

7 383 The new insights gained on the secondary aerosol formation and atmospheric
8 384 phenomena associated with air quality as a whole need to be disseminated from the
9 385 academia to the public and to the private sector. The academic experts need to keep
10 386 their knowledge and skills up-to-date and widen their knowledge base with horizontal
11 387 learning of the adjunct fields in science and technology. Atmospheric research
12 388 involves several fields of science, such as chemistry, physics, meteorology and Earth
13 389 system sciences, so deepening and widening the expertise is required. The horizontal
14 390 learning principle has been shown to be a good example of collaborative problem
15 391 solving and participatory action research [80]. The shift from discipline-tied
16 392 fundamental education towards a multi-disciplinarity is imperative for a successful
17 393 career in climate and global change science [81].
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33

34 394

35 395 In capacity building, we actually need to answer several questions: What are the target
36 396 groups? What knowledge needs to be transferred and what are the skills that each
37 397 target group needs? Concerning the comprehensive atmosphere earth system
38 398 measurements: Which kind of observation infrastructures is best to improve the air
39 399 quality in China? Concerning the effective knowledge transfer and innovative
40 400 thinking methods: What kind of knowledge transfer is needed for sustainable air
41 401 quality solutions?
42
43
44
45
46
47
48
49
50
51
52

53 402

54 403 Solution-oriented thinking, need for updating skills as well as knowledge of rapidly
55 404 changing air quality situation, are crucial. Reliable, research-based education that has
56
57
58
59
60

1
2
3 405 a holistic view on the whole big picture of causes and effects and their interactions
4
5 406 and feedbacks affecting air quality will support long lasting solutions. Also basic
6
7 407 understanding of the processes behind atmospheric phenomena is needed for building
8
9
10 408 a foundation for evaluating new information. Learning lasts a lifetime, which actually
11
12 409 is underlined by the fact that the university professors have pointed out that they
13
14 410 deepen their knowledge when lecturing to students.
15
16
17
18
19

20 412 **5. Future Outlook**

21
22
23 413
24
25
26 414 Atmospheric new particle formation contributes significantly to local, regional and
27
28 415 global aerosol number and CCN loads [e.g. 6]. Therefore, understanding of this
29
30 416 phenomenon is central to solving the secondary air pollution problem as a whole. The
31
32 417 following steps are needed in this process:
33
34
35

- 36 418 1) to perform long-term continuous, comprehensive observations on aerosol
37
38 419 precursors, oxidants, clusters, ions and aerosol particles together with proper
39
40 420 metadata and meteorological data. If needed, new Station for Measuring
41
42 421 Ecosystem – Atmosphere Relations II (SMEAR II, [82]) -type flagship stations
43
44 422 should be established, since they will help understanding the connections between
45
46 423 NPF and land surface – atmosphere interactions and feedbacks,
47
48
49 424 2) to establish calibration centers for mass spectrometers, PSMs and ion
50
51 425 spectrometers,
52
53 426 3) to organize joint data workshops for analyzing atmospheric data in proper,
54
55 427 comprehensive manner,
56
57
58
59
60

- 1
2
3 428 4) to ensure open data and metadata fluxes to other users, and
4
5 429 5) to organize joint paper writing workshops and publish the joint papers in peer-
6
7 430 reviewed journals.
8
9

10 431 It would be a big step forward to establish tight connections between different
11
12 432 Chinese research groups and support further deep collaborations in the future. The
13
14 433 second challenge is to establish open data policy and knowledge transfer at all levels.
15
16 434 The access to data is crucial to be able to answer research questions and to solve air
17
18 435 pollution problem(s). As a good sign, during the last years we have already seen
19
20 436 improvements regarding these issues. The third point is the capacity building,
21
22 437 including new infrastructures, data flows, databases etc. Furthermore, a new
23
24 438 generation of scientists needs to be educated to improve the knowledge base and
25
26 439 optimal use of infrastructures and data [86].
27
28
29
30
31
32
33

34 441 Understanding the formation of secondary pollutants is extremely important, since it
35
36 442 enables deep understanding of air pollutant dynamics crucial to air quality.
37
38 443 Improving air quality in China has several co-benefits, as it will lead to reduced
39
40 444 greenhouse-gas and black carbon emissions and concentrations, together with
41
42 445 improved fresh water quality and food supply. The cleaner air will decrease adverse
43
44 446 health effects caused by pollutants significantly [83, 84]. Efforts to prevent adverse
45
46 447 health effects must be well planned and should occur on multiple levels and places
47
48 448 simultaneously. Successful efforts will lead to significant gains in population health,
49
50 449 personal well-being and environmental quality as well as improving economy in
51
52 450 personal, local and national levels together with other significant co-benefits [85].
53
54
55
56
57
58
59
60

1
2
3 451 Reducing the use of fossil fuels does not only reduce emissions of air pollutants, but
4
5 452 also CO₂ and black carbon (BC), thereby decreasing radiative forcing in national and
6
7 453 global scales. Also, agricultural production and ecosystem services will benefit from
8
9
10 454 lowered pollutant levels. Healthier food will further improve peoples' health, and less
11
12 455 pollution damage improves yields of vegetables and crops. Better insulation of
13
14 456 buildings will lower the need for indoor heating, thus reducing emissions, but can also
15
16 457 reduce outdoor-indoor penetration of air pollutants. New technology in industry,
17
18 458 traffic and energy production will decrease emissions. The reduced pollution will
19
20
21 459 increase solar radiation in ground level and increase potential for solar energy.
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

460

461 Thus, tackling the air quality rapidly can lead to significant improvement on the
462 quality of life of the population as a whole and can lead to a positive feedback cycle,
463 which will encourage further progress towards cleaner environment.

464

465 **References**

466

- 467 1. IPCC. 2013. Climate Change 2013: The Physical Science Basis. Stocker T F,
468 Qin D, Plattner G K, Tignor M, Allen S K, Boschung J, Nauels A, Xia Y, Bex
469 V, Midgley P M, eds. Cambridge University Press, Cambridge, United
470 Kingdom and New York, NY, USA, 1535 pp,
471 doi:10.1017/CBO9781107415324

- 1
2
3 472 2. Hand J L, Malm W C. Review of aerosol mass scattering efficiencies from
4
5 473 ground-based measurements since 1990. 2007. *Journal of Geophysical*
6
7 474 *Research*, 112: D16203. doi:10.2029/2007JD008484
8
9
10 475 3. Lelieveld J, Evans J S, Fnais M, Giannadaki D, Pozzer A. The contribution of
11
12 476 outdoor pollution sources to premature mortality on a global scale. 2015.
13
14 477 *Nature*, 535: 367-371
15
16
17 478 4. Kulmala M, Nieminen T, Nikandrova A, Lehtipalo K, Manninen H E, Kajos
18
19 479 M K, Kolari P, Lauri A, Petäjä T, Krejci R, Hansson H-C, Swietlicki E,
20
21 480 Lindroth A, Christensen T R, Arneth A, Hari P, Bäck J, Vesala T, Kerminen
22
23 481 V-M. CO₂-induced terrestrial climate feedback mechanism: From carbon sink
24
25 482 to aerosol source and back. 2014a. *Boreal Environmental Research*, 19: suppl.
26
27 483 B, 122-131
28
29
30 484 5. Kulmala M, Petäjä T, Ehn M, Thornton J, Sipilä M, Worsnop D R, Kerminen
31
32 485 V-M. Chemistry of atmospheric nucleation: On the recent advances on
33
34 486 precursor characterization and atmospheric cluster composition in connection
35
36 487 with atmospheric new particle formation. 2014b. *Annual Review of Physical*
37
38 488 *Chemistry*, 65: 21-37
39
40
41 489 6. Kerminen V-M, Paramonov M, Anttila T, Riipinen I, Fountoukis C, Korhonen
42
43 490 H, Asmi E, Laakso L, Lihavainen H, Swietlicki E, Svenningsson B, Asmi A,
44
45 491 Pandis S N, Kulmala M, Petäjä T. Cloud condensation nuclei production
46
47 492 associated with atmospheric nucleation: a synthesis based on existing
48
49 493 literature and new results. 2012. *Atmospheric Chemistry and Physics*, 12:
50
51 494 12037-12059
52
53
54
55
56
57
58
59
60

- 1
2
3 495 7. Kulmala M, Kerminen V-M. On the formation and growth of atmospheric
4
5 496 nanoparticles. 2008. Atmospheric Research, 90: 132-150
6
7
8 497 8. Zhang R, Khalizov A, Wang L, Hu M, Xu W. Nucleation and growth of
9
10 498 nanoparticles in the atmosphere. 2012. Chemical Reviews, 112: 1957–2011
11
12
13 499 9. Weber R J, Marti J J, McMurry P H, Eisele F L, Tanner D J, Jefferson.
14
15 500 Measured atmospheric new particle formation rates: Implications for
16
17 501 nucleation mechanisms. 1996. Chemical Engineering Communication, 151:
18
19 502 53–64
20
21
22
23 503 10. Kulmala M, Lehtinen K E J, Laaksonen A. Cluster activation theory as an
24
25 504 explanation of the linear dependence between formation rate of 3 nm particles
26
27 505 and sulphuric acid concentration. 2006. Atmospheric Chemistry and Physics,
28
29 506 6: 787–793. doi:10.5194/acp-6-787-2006
30
31
32
33 507 11. Kulmala M, Toivonen A, Mäkelä J, Laaksonen A. Analysis of the growth of
34
35 508 nucleation mode particles observed in Boreal forest. 1998. Tellus B, 50: 449–
36
37 509 462
38
39
40 510 12. Almeida J, Schobesberger S, Kurten A, Ortega I K, Kupiainen-Määttä O,
41
42 511 Praplan A P, Adamov A, Amorim A, Bianchi F, Breitenlechner M, David A,
43
44 512 Dommen J, Donahue N M, Downard A, Dunne E, Duplissy J, Ehrhart S,
45
46 513 Flagan R C, Franchin A, Guida R, Hakala J, Hansel A, Heinritzi M, Henschel
47
48 514 H, Jokinen T, Junninen H, Kajos M, Kangasluoma J, Keskinen H, Kupc A,
49
50 515 Kurten T, Kvashin A N, Laaksonen A, Lehtipalo K, Leiminger M, Leppä J,
51
52 516 Loukonen V, Makhmutov V, Mathot S, McGrath M J, Nieminen T, Olenius T,
53
54 517 Onnela A, Petäjä T, Riccobono F, Riipinen I, Rissanen M, Rondo L,
55
56
57
58
59
60

- 1
2
3 518 Ruuskanen T, Santos F D, Sarnela N, Schallhart S, Schnitzhofer R, Seinfeld J
4
5 519 H, Simon M, Sipilä M, Stozhkov Y, Stratmann F, Tome A, Tröstl J,
6
7 520 Tsagkogeorgas G, Vaattovaara P, Viisanen Y, Virtanen A, Vrtala A, Wagner
8
9
10 521 P E, Weingartner E, Wex H, Williamson C, Wimmer D, Ye P L, Yli-Juuti T,
11
12 522 Carslaw K S, Kulmala M, Curtius J, Baltensperger U, Worsnop D R,
13
14 523 Vehkamäki H, Kirkby J. Molecular understanding of sulphuric acid-amine
15
16 524 particle nucleation in the atmosphere. *Nature*, 2013, 502: 359-363
- 17
18
19 525 13. Ehn M, Thornton J A, Kleist E, Sipilä M, Junninen H, Pullinen I, Springer M,
20
21 526 Rubach F, Tillmann R, Lee B, Lopez-Hifiker F, Andres S, Acir I H, Rissanen
22
23 527 M, Jokinen T, Schobesberger S, Kangasluoma J, Kontkanen J, Nieminen T,
24
25
26 528 Kurten T, Nielsen L B, Jorgensen S, Jaergaard H G, Canagaratna M, Dal Maso
27
28 529 M, Berndt T, Petäjä T, Wahner A, Kerminen V-M, Kulmala M, Worsnop D,
29
30 530 Wildt J, Mentel T F. A large source of low-volatility secondary organic
31
32 531 aerosol. 2014. *Nature*, 506: 476-479
- 33
34
35
36 532 14. Jokinen T, Berndt T, Makkonen R, Kerminen V-M, Junninen H, Paasonen P,
37
38 533 Stratmann F, Herrmann H, Guenther A, Worsnop D R, Kulmala M, Ehn M,
39
40 534 Sipilä M. Production of extremely low-volatile organic compounds from
41
42 535 biogenic emissions: measured yields and atmospheric implications. 2015.
43
44 536 *Proceedings of the National Academy of Sciences of the United States of*
45
46 537 *America*, 112: 7123-7128
- 47
48
49 538 15. McMurry P H, Friedlander S K. New particle formation in the presence of
50
51 539 aerosol. 1979. *Atmospheric Environment*, 13: 1635-1651
- 52
53
54
55
56
57
58
59
60

- 1
2
3 540 16. Kerminen V-M, Pirjola L, Kulmala M. How significantly does coagulation
4 scavenging limit atmospheric particle production? 2001. Journal of
5 541
6 Geophysical Research, 106: 24119-24126
7 542
8
9
10 543 17. Lehtinen K E J, Dal Maso M, Kulmala M, Kerminen V-M. Estimating
11 nucleation rates from apparent particle formation rates and vice-versa: Revised
12 544
13 formulation of the Kerminen-Kulmala equation. 2007. Journal of Aerosol
14 545
15 Science, 38: 988-994
16 546
17
18
19
20 547 18. Kulmala M, Kontkanen J, Junninen H, Lehtipalo K, Manninen H E, Nieminen
21 T, Petäjä T, Sipilä M, Schobesberger S, Rantala P, Franchin A, Jokinen T,
22 548
23 Järvinen E, Äijälä M, Kangasluoma J, Hakala J, Aalto P P, Paasonen P,
24 549
25 Mikkilä J, Vanhanen J, Aalto J, Hakola H, Makkonen U, Ruuskanen T,
26 550
27 Mauldin III R L, Duplissy J, Vehkamäki H, Bäck J, Kortelainen A, Riipinen I,
28 551
29 Kurten T, Johnston M V, Smith J N, Ehn M, Mentel T F, Lehtinen K E J,
30 552
31 Laaksonen A, Kerminen V-M, Worsnop D R. Direct observations of
32 553
33 atmospheric aerosol nucleation. 2013. Science, 339: 943-946
34 554
35
36
37
38 555 19. Vehkamäki H, Riipinen I. Thermodynamics and kinetics of atmospheric
39 aerosol particle formation and growth. 2012. Chemistry Society Review, 41
40 556
41
42
43 557 20. Buonanno G, Marks G B, Morawska L. Health effects of daily airborne
44 particle dose in children: Direct association between personal dose and
45 558
46 respiratory health effects. Environmental Pollution, 2013, 180: 246-250
47 559
48
49
50 560 21. Kulmala M, Kerminen V-M, Anttila T, Laaksonen A, O'Dowd C D. Organic
51 aerosol formation via sulphate cluster activation. 2004. Journal of Geophysical
52 561
53 Research, 109 (D4). doi: 10.1029/2003JD003961
54 562
55
56
57
58
59
60

- 1
2
3 563 22. Rosenfeld D, Sherwood S, Wood R, Donner L. Climate effects of aerosol-
4
5 564 cloud interactions. 2014. *Science*, 343: 379-380
6
7
8 565 23. Qu W J, Wang J, Zhang X Y, Wang D, Sheng L F. Influence of relative
9
10 566 humidity on aerosol composition: Impacts on light extinction and visibility
11
12 567 impairment at two sites in coastal area of China. 2015. *Atmospheric Research*,
13
14 568 153: 500-511
15
16
17
18 569 24. Nemmar A, Hoet P H, Vanquickenborne B, Dinsdate D, Thomeer M,
19
20 570 Hoylaerts M F, Valbilloen H, Mortelmans L, Nemery B. Passage of inhaled
21
22 571 particles into the blood circulation in humans. 2002. *Circulation*, 105: 411-414
23
24
25 572 25. Guo S, Hua M, Zamorab M L, Peng J F, Shang D J, Zheng J, Du Z F, Wu Z J,
26
27 573 Shao M, Zeng L M, Molinac M J, Zhang R Y. Elucidating severe urban haze
28
29 574 formation in China. 2014. *Proceedings of the National Academy of Sciences*
30
31 575 *of the United States of America*, 111: 17373-17378
32
33
34
35 576 26. Huang R J, Zhang Y L, Bozzetti C, Ho K F, Cao J J, Han Y M, Daellenbach
36
37 577 K R, Slowik J G, Platt S M, Canonaco F, Zotter P, Wolf R, Pieber S M,
38
39 578 Bruns E A, Crippa M, Ciarelli G, Piazzalunga A, Schwikowski M,
40
41 579 Abbaszade G, Schnelle-Kreis J, Zimmermann R, An Z, Szidat S,
42
43 580 Baltensperger U, El Haddad I, Prévôt A S H. High secondary aerosol
44
45 581 contribution to particulate pollution during haze events in China. 2014. *Nature*
46
47 582 514: 218-222
48
49
50
51 583 27. Wu Z J, Hu M, Liu S, Wehner B, Bauer S, Maßling A, Wiedensohler A,
52
53 584 Petäjä T, Dal Maso M, Kulmala M. New particle formation in Beijing, China:

- 1
2
3 585 Statistical analysis of a 1-year data set. 2007. Journal of Geophysical
4
5 586 Research: Atmospheres, 112: D09209. doi 10.1029/2006jd007406
6
7
8 587 28. Xiao S, Wang M Y, Yao L, Kulmala M, Zhou B, Yang X, Chen J M, Wang D
9
10 588 F, Fu Q Y, Worsnop D R, Wang L. Strong atmospheric new particle formation
11
12 589 in winter in urban Shanghai, China. 2015. Atmospheric Chemistry and
13
14 590 Physics, 15: 1769-1781
15
16
17
18 591 29. Nie W, Ding A, Wang T, Kerminen V-M, George C, Xue L, Wang W, Zhang
19
20 592 Q, Petäjä T, Qi X, Wang X, Yang X, Fu C, Kulmala M. Polluted dust
21
22 593 promotes new particle formation and growth. 2014. Scientific Reports, 4:
23
24 594 6634, doi:10.1038/srep06634
25
26
27
28 595 30. Xie Y, Ding A, Nie W, Mao H, Qi X, Huang X, Xu Z, Kerminen V-M, Petäjä
29
30 596 T, Chi X, Virkkula A, Boy M, Xue L, Guo J, Sun J, Yang X, Kulmala M, Fu
31
32 597 C. Enhanced sulfate formation by nitrogen dioxide: Implications from in situ
33
34 598 observations at the SORPES station. 2015. Journal of Geophysical Research:
35
36 599 Atmospheres, 120, 24: 12679–12694
37
38
39 600 31. Kulmala M. China's choking cocktail. 2015. Nature, 526: 497-499
40
41
42 601 32. Fiore A M, Naik V, Spracklen D V, Steiner A, Unger N, Prather M, Bergmann
43
44 602 D, Cameron-Smith P J, Cionni I, Collins W J, Dalsoren S, Eyring V, Folberth
45
46 603 G A, Ginoux P, Horowitz L W, Josse B, Lamarque J F, MacKenzie I A,
47
48 604 Nagashima T, O'Connor F M, Righi M, Rumbold S T, Shindell D T, Skeie R
49
50 605 B, Sudo K, Szopa S, Takemura T, Zeng G. Global air quality and climate.
51
52 606 2012. Chemical Society Reviews, 41: 6663-6683. doi:10.1039/c2cs35095e
53
54
55
56
57
58
59
60

- 1
2
3 607 33. Fuzzi S, Baltensperger U, Carslaw K, Decesari S, Denier van der Gon H,
4
5 608 Facchini M C, Fowler D, Koren I, Langford B, Lohmann U, Nemitz E, Pandis
6
7 609 S, Riipinen I, Rudich Y, Schaap M, Slowik J G, Spracklen D V, Vignati E,
8
9 610 Wild M, Williams M, Gilardoni S. Particulate matter, air quality and climate:
10
11 611 lessons learned and future needs. 2015. Atmospheric Chemistry and Physics,
12
13 612 15: 8217-8299
- 14
15
16
17 613 34. Hari P, Petäjä T, Bäck J, Kerminen V-M, Lappalainen H K, Vihma T, Laurila
18
19 614 T, Viisanen Y, Vesala T, Kulmala M. Conceptual design of a measurement
20
21 615 network of the global change. 2016. Atmospheric Chemistry and Physics, 16:
22
23 616 1017-1028
- 24
25
26
27 617 35. Ding A, Fu C, Yang X, Sun J, Zheng L, Xie Y, Herrmann E, Nie W, Petäjä T,
28
29 618 Kerminen V-M, Kulmala M. Ozone and fine particle in the western Yangtze
30
31 619 river delta: an overview of 1 yr data at the SORPES station. 2013a.
32
33 620 Atmospheric Chemistry and Physics, 13: 5813-5830
- 34
35
36
37 621 36. Ding A J, Fu C B, Yang X Q, Sun J N, Petäjä T, Kerminen V-M, Wang T, Xie
38
39 622 Y N, Herrmann E, Zheng L F, Nie W, Wei L W, Kulmala M. Intense
40
41 623 atmospheric pollution modifies weather: a case of mixed biomass burning with
42
43 624 fossil fuel combustion pollution in the eastern China. 2013b. Atmospheric
44
45 625 Chemistry and Physics, 13: 10545-10554
- 46
47
48 626 37. Petäjä T, Järvi L, Kerminen V-M, Ding A, Sun J, Nie W, Kujansuu J,
49
50 627 Virkkula A, Yang X, Fu C, Zilitinkevich S, Kulmala M. Air pollution: a new
51
52 628 wall of China. 2016. Scientific Reports

- 1
2
3 629 38. Lin P, Hu M, Wu Z, Niu Y, Zhu T. Marine aerosol size distributions in the
4
5 630 springtime over China adjacent seas. 2007 Atmospheric Environment, 41:
6
7 631 6784-6796, doi10.1016/j.atmosenv.2007.04.045
8
9
10 632 39. Liu S, Hu M, Wu Z J, Wehner B, Wiedensohler A, Cheng Y F. Aerosol
11
12 633 number size distribution and new particle formation at a rural/coastal site in
13
14 634 Pearl River Delta (PRD) of China. 2008. Atmospheric Environment, 42: 6275-
15
16 635 6283, doi10.1016/j.atmosenv.2008.01.063
17
18
19
20 636 40. Gong Y G, Hu M, Cheng Y, Su H, Yue D, Liu F, Wiedensohler A, Wang Z,
21
22 637 Kalesse H, Liu S, Wu Z, Xiao K, Mi P, Zhang Y. Competition of coagulation
23
24 638 sink and source rate: New particle formation in the Pearl River Delta of China.
25
26 639 2010. Atmospheric Environment, 44: 3278-3285. doi:
27
28 640 10.1016/j.atmosenv.2010.05.049
29
30
31
32 641 41. Yue D L, Hu M, Zhang R Y, Wang Z B, Zheng J, Wu Z J, Wiedensohler A,
33
34 642 He L Y, Huang X F, Zhu T. The roles of sulfuric acid in new particle
35
36 643 formation and growth in the mega-city of Beijing. 2010. Atmospheric
37
38 644 Chemistry and Physics, 10: 4953-4960
39
40
41
42 645 42. Wu Z J, Hu M, Yue D L, Liu S, Wehner B, Wiedensohler A. Evolution of
43
44 646 particle number size distribution in an urban atmosphere during episodes of
45
46 647 heavy pollution and new particle formation. 2011. Science China Earth
47
48 648 Science, 54: 1772-1778
49
50
51
52 649 43. Yue D L, Hu M, Zhang R Y, Wu Z J, Su H, Wang Z B, Peng J F, He L Y,
53
54 650 Huang X F, Gong Y G, Wiedensohler A. Potential contribution of new particle
55
56
57
58
59
60

- 1
2
3 651 formation to cloud condensation nuclei in Beijing. 2011. Atmospheric
4
5 652 Environment, 45: 6070-6077
6
7
8 653 44. Wang Z B, Hu M, Mogensen D, Yue D L, Zheng J, Zhang R Y, Liu Y, Yuan
9
10 654 B, Li X, Shao M, Zhou L, Wu Z J, Wiedensohler A, Boy M. The simulations
11
12 655 of sulfuric acid concentration and new particle formation in an urban
13
14 656 atmosphere in China. 2013a. Atmospheric Chemistry and Physics, 13: 11157-
15
16 657 11167. 10.5194/acp-13-11157-2013
17
18
19
20 658 45. Yue D L, Hu M, Wang Z B, Wen M T, Guo S, Zhong L J, Wiedensohler A,
21
22 659 Zhang Y H. Comparison of particle number size distributions and new
23
24 660 particle formation between the urban and rural sites in the PRD region, China.
25
26 661 2013. Atmospheric Environment, 76: 181-188
27
28
29
30 662 46. Peng J F, Hu M, Wang Z B, Huang X F, Kumar P, Wu Z J, Guo S, Yue D L,
31
32 663 Shang D J, Zheng Z, He L Y. Submicron aerosols at thirteen diversified sites
33
34 664 in China: size distribution, new particle formation and corresponding
35
36 665 contribution to cloud condensation nuclei production. 2014. Atmospheric
37
38 666 Chemistry and Physics, 14: 10249-10265, 10.5194/acp-14-10249-2014
39
40
41
42 667 47. Qi X, Ding A J, Nie W, Petäjä T, Kerminen V-M, Herrmann E, Xie Y N,
43
44 668 Zheng L F, Manninen H, Aalto P, Sun J N, Xu Z N, Chi X G, Huang X, Boy
45
46 669 M, Virkkula A, Yang X Q, Fu C B, Kulmala M. Aerosol size distribution and
47
48 670 new particle formation in western Yangtze River Delta of China: two-year
49
50 671 measurement at the SORPES station. 2015. Atmospheric Chemistry and
51
52 672 Physics Discussions, 15: 12491-12537
53
54
55
56
57
58
59
60

- 1
2
3 673 48. Wehner B, Wiedensohler A, Tuch T M, Wu Z J, Hu M, Slanina J, Kiang C S.
4
5 674 Variability of the aerosol number size distribution in Beijing, China: New
6
7 675 particle formation, dust storms, and high continental background. 2004.
8
9 676 Geophysical Research Letters, 31: L22108
- 10
11
12 677 49. Wang Z B, Hu M, Wu Z J, Yue D L, He L Y, Huang X F, Liu X G,
13
14 678 Wiedensohler A. Long-term measurements of particle number size
15
16 679 distributions and the relationships with air mass history and source
17
18 680 apportionment in the summer of Beijing. 2013d. Atmospheric Chemistry and
19
20 681 Physics, 13: 10159-10170
- 21
22
23
24 682 50. Wang Z B, Hu M, Wu Z J, Yue D L. Research on the Formation Mechanisms
25
26 683 of New Particles in the Atmosphere. 2013c. Acta Chimica Sinica, 71: 519-527
- 27
28
29
30 684 51. Yue D L, Hu M, Wu Z J, Wang Z B, Guo S, Wehner B, Nowak A, Achtert P,
31
32 685 Wiedensohler A, Jung J S, Kim Y J, Liu S C. Characteristics of aerosol size
33
34 686 distributions and new particle formation in the summer in Beijing. 2009.
35
36 687 Journal of Geophysical Research: Atmospheres, 114, D00G12
- 37
38
39 688 52. Wang Z B, Hu M, Yue D L, Zheng J, Zhang R Y, Wiedensohler A, Wu Z J,
40
41 689 Nieminen T, Boy M. Evaluation on the role of sulfuric acid in the mechanisms
42
43 690 of new particle formation for Beijing case. 2011. Atmospheric Chemistry and
44
45 691 Physics, 11: 12663-12671. 10.5194/acp-11-12663-2011
- 46
47
48
49 692 53. Wang Z B, Hu M, Pei X Y, Zhang R Y, Paasonen P, Zheng J, Yue D L, Wu Z
50
51 693 J, Boy M, Wiedensohler A. Connection of organics to atmospheric new
52
53 694 particle formation and growth at an urban site of Beijing. 2015. Atmospheric
54
55 695 Environment, 103: 7-17

- 1
2
3 696 54. Kulmala M, Dal Maso M, Mäkelä J M, Pirjola L, Väkevä M, Aalto P,
4
5 697 Miikkulainen P, Hämeri K, O'Dowd C D. On the formation, growth and
6
7 698 composition of nucleation mode particles. 2001. *Tellus*, 53B: 479-480
8
9
10 699 55. Kulmala M, Petäjä T, Mönkkönen P, Koponen I K, Dal Maso M, Aalto P P,
11
12 700 Lehtinen K E J, Kerminen V-M. On the growth of nucleation mode particles:
13
14 701 source rates of condensable vapor in polluted and clean environments. 2005.
15
16 702 *Atmospheric Chemistry and Physics*, 5: 409-416
17
18
19
20 703 56. Wang Z B, Hu M, Sun J Y, Wu Z J, Yue D L, Shen X J, Zhang Y M, Pei X Y,
21
22 704 Cheng Y F, Wiedensohler A. Characteristics of regional new particle
23
24 705 formation in urban and regional background environments in the North China
25
26 706 Plain. 2013b. *Atmospheric Chemistry and Physics*, 13: 12495-12506.
27
28 707 10.5194/acp-13-12495-2013
29
30
31
32 708 57. Wang Z B, Hu M, Yue D L, He L Y, Huang X F, Yang Q, Zheng J, Zhang R
33
34 709 Y, Zhang Y H. New particle formation in the presence of a strong biomass
35
36 710 burning episode at a downwind rural site in PRD, China. 2013e. *Tellus B*, 65:
37
38 711 19965
39
40
41
42 712 58. Dupart Y, King S M, Nekat B, Nowak A, Wiedensohler A, Herrmann H,
43
44 713 David G, Thomas B, Miffre A, Rairoux P, D'Anna B, George C. Mineral dust
45
46 714 photochemistry induces nucleation events in the presence of SO₂. 2012.
47
48 715 *Proceedings of the National Academy of Sciences*, 109: 20842-20847. doi:
49
50 716 10.1073/pnas.1212297109
51
52
53
54 717 59. Nie W, Ding A J, Xie Y N, Xu Z, Mao H, Kerminen V M, Zheng L F, Qi X
55
56 718 M, Huang X, Yang X Q, Sun J N, Herrmann E, Petäjä T, Kulmala M, Fu C B.

- 1
2
3 719 Influence of biomass burning plumes on HONO chemistry in eastern China.
4
5 720 2015. *Atmospheric Chemistry and Physics*, 15: 1147-1159
6
7
8 721 60. He H, Wang Y, Ma Q, Ma J, Chu B, Ji D, Tang G, Liu C, Zhang H, Hao J.
9
10 722 Mineral dust and NO_x promote the conversion of SO₂ to sulfate in heavy
11
12 723 pollution days. 2014. *Scientific Reports*, 4: 4172
13
14
15 724 61. Vanhanen J, Mikkilä J, Lehtipalo K, Sipilä M, Manninen H E, Siivola E,
16
17 725 Petäjä T, Kulmala M. Particle size magnifier for nano-CN Detection. 2011.
18
19 726 *Aerosol Science and Technology*, 45: 533-542
20
21
22
23 727 62. Kulmala M, Riipinen I, Sipilä M, Manninen H E, Petäjä T, Junninen H, Dal
24
25 728 Maso M, Mordas G, Mirme A, Vana M, Hirsikko A, Laakso L, Harrison R M,
26
27 729 Hanson I, Leung C, Lehtinen K E J, Kerminen V-M. Towards direct
28
29 730 measurement of atmospheric nucleation. 2007. *Science*, 318: 89-92
30
31
32
33 731 63. Tammet H. Symmetric Inclined Grid Mobility Analyzer for the Measurement
34
35 732 of Charged Clusters and Fine Nanoparticles in Atmospheric Air. 2011.
36
37 733 *Aerosol Science and Technology*, 45: 468-479
38
39
40 734 64. Kulmala M, Petäjä T, Nieminen T, Sipilä M, Manninen H E, Lehtipalo K, Dal
41
42 735 Maso M, Aalto P P, Junninen H, Paasonen P, Riipinen I, Lehtinen K E J,
43
44 736 Laaksonen A, Kerminen V-M. Measurement of the nucleation of atmospheric
45
46 737 aerosol particles. 2012. *Nature Protocols*, 7: 1651-1667
47
48
49
50 738 65. Lehtipalo K, Leppä J, Kontkanen J, Kangasluoma J, Franchin A, Wimmer D,
51
52 739 Schobesberger S, Junninen H, Petäjä T, Sipilä M, Mikkilä J, Vanhanen J,
53
54 740 Worsnop D R, Kulmala M. Methods for determining particle size distribution
55
56
57
58
59
60

- 1
2
3 741 and growth rates between 1 and 3 nm using the Particle Size Magnifier. 2014.
4
5 742 Boreal Environmental Research, 19: B, 215-236
6
7
8 743 66. Gagné S, Nieminen T, Kurtén T, Manninen H E, Petäjä T, Laakso L,
9
10 744 Kerminen V-M, Boy M, Kulmala M. Factors influencing the contribution of
11
12 745 ion-induced nucleation in a boreal forest, Finland. 2010. Atmospheric
13
14 746 Chemistry and Physics, 10: 3743-3757
15
16
17
18 747 67. Kulmala M, Riipinen I, Nieminen T, Hulkkonen M, Sogacheva L, Manninen
19
20 748 H E, Paasonen P, Petäjä T, Dal Maso M, Aalto P P, Viljanen A, Usoskin I,
21
22 749 Vainio R, Mirme S, Mirme A, Minikin A, Petzold A, Härrak U, Plaß-Dülmer
23
24 750 C, Birmili, Kerminen V-M. Atmospheric data over a solar cycle: no
25
26 751 connection between galactic cosmic rays and new particle formation. 2010.
27
28 752 Atmospheric Chemistry and Physics, 10: 1885-1898
29
30
31
32 753 68. Junninen H, Ehn M, Petäjä T, Luosujärvi L, Kotiaho T, Kostianen R, Rohner
33
34 754 U, Gonin M, Fuhrer K, Kulmala M, Worsnop D R. A high-resolution mass
35
36 755 spectrometer to measure atmospheric ion composition. 2010. Atmospheric
37
38 756 Measurement Techniques, 3: 1039-1053. doi:10.5194/amt-3-1039-2010
39
40
41
42 757 69. Jokinen T, Sipilä M, Junninen H, Ehn M, Lönn G, Hakala J, Petäjä T, Mauldin
43
44 758 III R L, Kulmala M, Worsnop D R. Atmospheric sulfuric acid and neutral
45
46 759 cluster measurements using CI-API-TOF. 2012. Atmospheric Chemistry and
47
48 760 Physics, 12: 4117-4125
49
50
51 761 70. Petäjä T, Mauldin III R L, Kosciuch E, McGrath J, Nieminen T, Paasonen P,
52
53 762 Boy M, Adamov A, Kotiaho T, Kulmala M. Sulfuric acid and OH
54
55
56
57
58
59
60

- 1
2
3 763 concentrations in a boreal forest site. 2009. Atmospheric Chemistry and
4
5 764 Physics, 9: 7435-7448
6
7
8 765 71. Sipilä, M, Berndt T, Petäjä T, Brus D, Vanhanen J, Stratmann F, Patokoski J,
9
10 766 Mauldin III R L, Hyvärinen A P, Lihavainen H, Kulmala M. The role of
11
12 767 sulfuric acid in atmospheric nucleation. 2010. Science, 327: 1243-1246
13
14
15 768 72. Kirkby J, Curtius J, Almeida J, Dunne E, Duplissy J, Ehrhart S, Franchin A,
16
17 769 Gagné S, Ickes L, Kürten A, Kupc A, Metzger A, Riccobono F, Rondo L,
18
19 770 Schobesberger S, Tsagkogeorgas G, Wimmer D, Amorim A, Bianchi F,
20
21 771 Breitenlechner M, David A, Dommen J, Downard A, Ehn M, Flagan R C,
22
23 772 Haider S, Hansel A, Hauser D, Jud W, Junninen H, Kreissl F, Kvashin A,
24
25 773 Laaksonen A, Lehtipalo K, Lima J, Lovejoy E R, Makhmutov V, Mathot S,
26
27 774 Mikkilä J, Minginette P, Mogo S, Nieminen T, Onnela A, Pereira P, Petäjä T,
28
29 775 Schnitzhofer R, Seinfeld J H, Sipilä M, Stozhkov Y, Stratmann F, Tomé A,
30
31 776 Vanhanen J, Viisanen Y, Vrtala A, Wagner P E, Walther H, Weingartner E,
32
33 777 Wex H, Winkler P M, Carslaw K S, Worsnop D R, Baltensperger U, Kulmala
34
35 778 M. The role of sulfuric acid, ammonia and galactic cosmic rays in atmospheric
36
37 779 aerosol nucleation. 2011. Nature, 476: 429-433
38
39
40
41
42
43 780 73. Petäjä T, Sipilä M, Paasonen P, Nieminen T, Kurtén T, Ortega I K, Stratmann
44
45 781 F, Vehkamäki H, Berndt T, Kulmala M. Experimental observation of strongly
46
47 782 bound dimers of sulphuric acid: application to nucleation in the atmosphere.
48
49 783 2011. Physical Review Letters, 106: 228302
50
51
52
53 784 74. Riccobono F, Schobesberger S, Scott C E, Dommen J, Ortega I K, Rondo L,
54
55 785 Almeida J, Amorim A, Bianchi F, Breitenlechner M, David A, Downard A,

- 1
2
3 786 Dunne E M, Duplissy J, Ehrhart S, Flagan R C, Franchin A, Hansel A,
4
5 787 Junninen H, Kajos M, Keskinen H, Kupc A, Kürten A, Kvashin A N,
6
7 788 Laaksonen A, Lehtipalo K, Makhmutov V, Mathot S, Nieminen T, Onnela A,
8
9 789 Petäjä T, Praplan AP, Santos F D, Schallhart S, Seinfeld J H, Sipilä M,
10
11 790 Spracklen D V, Stozhkov Y, Stratmann F, Tomé A, Tsagkogeorgas G,
12
13 791 Vaattovaara P, Viisanen Y, Vrtala A, Wagner P E, Weingartner E, Wex H,
14
15 792 Wimmer D, Carslaw K S, Curtius J, Donahue N M, Kirkby J, Kulmala M,
16
17 793 Worsnop D R, Baltensperger U. Oxidation products of biogenic emissions
18
19 794 contribute to nucleation of atmospheric particles. 2014. *Science*, 344: 717-721
20
21
22
23
24 795 75. Mauldin III R L, Berndt T, Sipilä M, Paasonen P, Petäjä T, Kim S, Kurtén T,
25
26 796 Stratmann F, Kerminen V-M, Kulmala M. New atmospherically relevant
27
28 797 oxidant. 2012. *Nature*, 488: 193-197
29
30
31 798 76. Taipale R, Sarnela N, Rissanen M, Junninen H, Rantala P, Korhonen F,
32
33 799 Siivola E, Berndt T, Kulmala M, Mauldin III R L, Petäjä T, Sipilä M. New
34
35 800 instrument for measuring atmospheric concentrations of non-OH oxidants of
36
37 801 SO₂. 2014. *Boreal Environment Research*, 19: B, 55-70
38
39
40
41 802 77. Mauldin III R L, Rissanen M P, Petäjä T, Kulmala M. Furthering information
42
43 803 from OH and HO₂+RO₂ observations using a high resolution time of flight
44
45 804 mass spectrometer. 2016. *Atmospheric Measurement Techniques Discussions*,
46
47 805 doi:10.5194/amt-2015-398
48
49
50 806 78. Wiedensohler A, Birmili W, Nowak A, Sonntag A, Weinhold K, Merkel M,
51
52 807 Wehner B, Tuch T, Pfeifer S, Fiebig M, Fjaraa A M, Asmi E, Sellegri K,
53
54 808 Depuy R, Venzac H, Villani P, Laj P, Aalto P, Ogren J A, Swietlicki E,

- 1
2
3 809 Williams P, Roldin P, Quincey P, Huglin C, Fierz-Schmidhauser R, Gysel M,
4
5 810 Weingartner E, Riccobono F, Santos S, Gruning C, Faloon K, Beddows D,
6
7 811 Harrison R, Monahan C, Jennings S G, O'Dowd C D, Marinoni A, Horn H-G,
8
9
10 812 Keck J, Jiang J, Scheckman J, McMurry P H, Deng Z, Zhao C S, Moerman
11
12 813 M, Henzing B, de Leeuw G, Loschau G, Bastian S. Mobility particle size
13
14 814 spectrometers: harmonization of technical standards and data structure to
15
16 815 facilitate high quality long-term observations of atmospheric particle number
17
18 816 size distributions. 2012. *Atmospheric Measurement Technology*, 5: 657–685
- 20
21
22 817 79. Crippa M, Canonaco F, Lanz V A, Äijälä M, Allan J D, Carbone S, Capes G,
23
24 818 Ceburnis D, Dall'Osto M, Day D A, DeCarlo P F, Ehn M, Eriksson A, Freney
25
26 819 E, Hildebrandt Ruiz L, Hillamo R, Jimenez J L, Junninen H, Kiendler-Scharr
27
28 820 A, Kortelainen A-M, Kulmala M, Laaksonen A, Mensah A A, Mohr C,
29
30 821 Nemitz E, O'Dowd C, Ovadnevaite J, Pandis S N, Petäjä T, Poulain L,
31
32 822 Saarikoski S, Sellegri K, Swietlicki E, Tiitta P, Worsnop D R, Baltensperger
33
34 823 U, Prévôt A S H. Organic aerosol components derived from 25 AMS data sets
35
36 824 across Europe using a consistent ME-2 based source apportionment approach.
37
38 825 2014. *Atmospheric Chemistry and Physics*, 14: 6159-6176
- 40
41
42 826 80. Hennessy S, Murphy P. The Potential for Collaborative Problem Solving in
43
44 827 Design and Technology. 1999. *International Journal of Technology and*
45
46 828 *Design Education* 9: 1, 1–36
- 48
49
50 829 81. Nordic Climate Change Research. 2009: NordForsk Policy Briefs 2009-8.
51
52 830 Mandag Morgen, 2009

- 1
2
3 831 82. Hari P, Kulmala M. Station for Measuring Ecosystem – Atmosphere Relations
4
5 832 (SMEAR II). 2005. *Boreal Environment Research*, 10: 315-322
6
7
8 833 83. Zhang J, Mauzerall D L, Zhu T, Liang S, Ezzati M, Remais J V.
9
10 834 Environmental health in China: progress towards clean air and safe water.
11
12 835 2010. *Lancet*, 375: 9720
13
14
15
16 836 84. Tang D, Wang C, Nie J, Chen R, Niu Q, Kan H, Chen B, Perera F. Health
17
18 837 benefits of improving air quality in Taiyuan, China. 2014. *Environment*
19
20 838 *International*, 73: 235-242. <http://dx.doi.org/10.1016/j.envint.2014.07.016>
21
22
23 839 85. Haines A, McMichael A J, Smith K R, Roberts I, Woodcock J, Markandya A,
24
25 840 Armstrong B G, Campbell-Lendrum D, Dangour A D, Davies M, Bruce N,
26
27 841 Tonne C, Barrett M, Wilkinson P. Public health benefits of strategies to reduce
28
29 842 greenhouse-gas emissions: overview and implications for policy makers. 2009.
30
31 843 *Lancet*, 374: 2104
32
33
34
35 844 86. Lappalainen H K, Kerminen V M, Petäjä T, Kurten T, Baklanov A, Shvidenko
36
37 845 A, Bäck J, Vihma T, Alekseychik P, Arnold S, Arshinov M, Asmi E, Belan B,
38
39 846 Bobylev L, Chalov S, Cheng Y, Chubarova N, de Leeuw G, Ding A,
40
41 847 Dobrolyubov S, Dubtsov S, Dyukarev E, Elansky N, Eleftheriadis K, Esau I,
42
43 848 Filatov N, Flint M, Fu C, Glezer O, Gliko A, Heimann M, Holtslag A M,
44
45 849 Hörrak U, Janhunen J, Juhola S, Järvi L, Järvinen H, Kanukhina A,
46
47 850 Konstantinov P, Kotlyakov V, Kieloaho A-J, Komarov A, Kujansuu J,
48
49 851 Kukkonen I, Kyrö E, Laaksonen A, Laurila T, Lihavainen H, Lisitzin A,
50
51 852 Mahura A, Makshtas A, Mareev E, Mazon S, Matishov D, Melnikov V,
52
53 853 Mikhailov E, Moisseev D, Nigmatulin R, Noe S M, Ojala A, Pihlatie M,
54
55
56
57
58
59
60

1
2
3 854 Popovicheva O, Pumpanen J, Regerand T, Repina I, Shcherbinin A,
4
5 855 Shevchenko V, Sipilä M, Skorokhod A, Spracklen D V, Su H, Subetto D, Sun
6
7 856 J, Terzhevik A, Timofeyev Y, Troitskaya Y, Tynkkynen V-P, Kharuk V I,
8
9
10 857 Zaytseva N, Zhang J, Viisanen Y, Vesala T, Hari P, Hansson H-C, Matvienko
11
12 858 G, Kasimov N, Guo H, Bondur V, Zilitinkevich S, Kulmala M. Pan-Eurasian
13
14 859 Experiment (PEEX): Towards holistic understanding of the feedbacks and
15
16 860 interactions in the land - atmosphere - ocean- society continuum in the
17
18 861 Northern Eurasian region. 2016. Atmospheric Chemistry and Physics - PEEX
19
20
21 862 Special Issue, submitted
22
23
24 863
25
26
27 864 **Acknowledgements**
28
29
30 865 The work in this manuscript is supported by Academy of Finland via Center of
31
32 866 Excellence in Atmospheric Sciences (project no. 272041) and the Finnish Funding
33
34 867 Agency for Technology and Innovation TEKES via Beautiful Beijing project
35
36 868 (3667/31/2013) and European Research Council Advanced Grant (ATMNUCLE,
37
38 869 227463) and InGOS DEFROST and CRAICC (no 26060) and Nordforsk CRAICC-
39
40 870 PEEX (amendment to contact 26060) funded by Nordforsk. The SORPES station was
41
42 871 supported by Nanjing University and the Collaborative Innovation Center of Climate
43
44 872 Change in Jiangsu Province, China. Part of Aijun Ding's work was supported by the
45
46 873 excellent young scientist fund of National Natural Science Foundation of China
47
48 874 (D0512/ 41422504).
49
50
51
52
53 875
54
55
56 876 X

877

For Review Only

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

1
2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26
27
28
29
30
31
32
33
34
35
36
37
38
39
40
41
42
43
44
45
46
47
48
49
50
51
52
53
54
55
56
57
58
59
60

Research highlights

1) Formation of new atmospheric aerosol particles is a global phenomenon that has been observed to take place in even heavily-polluted environments. A holistic scientific understanding on the atmospheric phenomena associated with air quality as a whole, as well as on the connection between air quality and climate, is lacking at the moment.

2) In China, new particle production has been observed at very high pollution levels (condensation sink about 0.1 s^{-1}) in several megacities. With a network of observation stations, we will be able to understand the interactions and feedbacks associated with the urban pollution mixture, and ultimately, be ready to make targeted strategies for the pollution control.

3) This paper summarizes the recent advances in studying secondary new aerosol formation in China to show how increased process-level understanding will help us to understand air quality-climate-weather interactions and how the feedbacks and interactions affect the air quality in highly-polluted environments such as those frequently encountered in Chinese megacities.

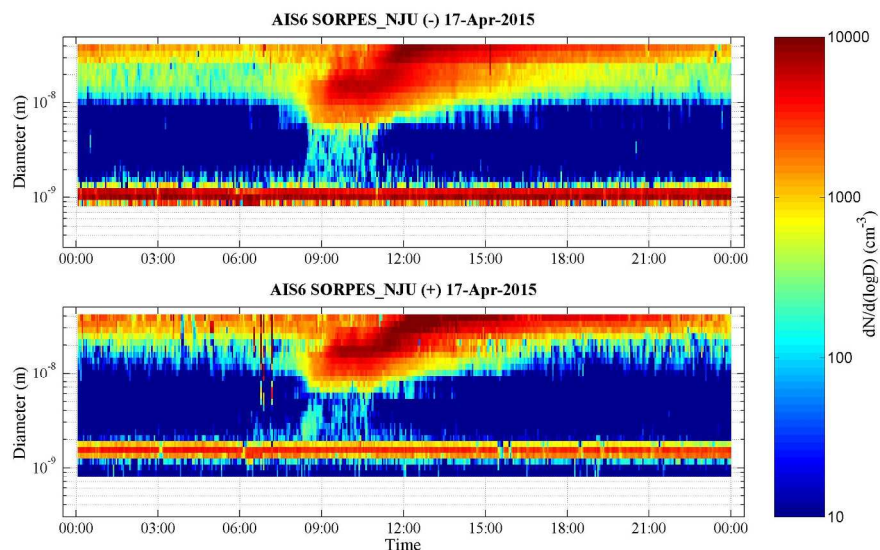


Figure 1. A typical nucleation event measured using Air Ion Spectrometer (AIS) at the SORPES station, Nanjing, in China. The background cluster ions are seen in both negative and positive ion modes in the sub-2 nm size range. Negative ion clusters are smaller than positive ones. The new particle formation is seen in both polarities starting at around 8.30 am. Here J_6 is $1.8 \text{ cm}^{-3}\text{s}^{-1}$ and $\text{GR}(6-30 \text{ nm})$ is 6.6 nm/hr .
871x523mm (72 x 72 DPI)